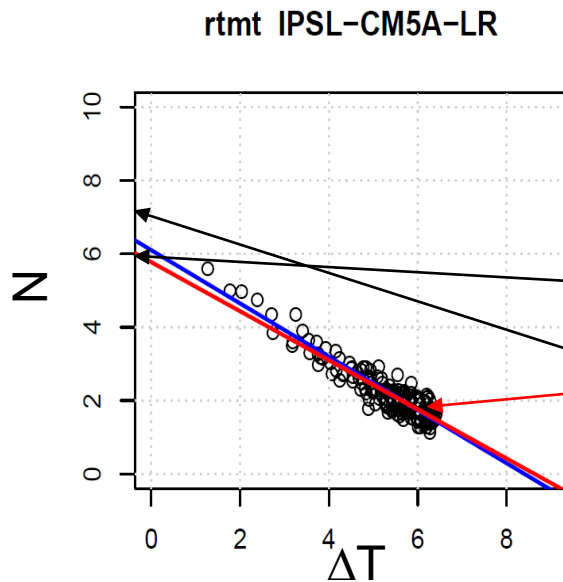


# On the interpretation of inter-model spread in CMIP5 climate sensitivity estimates

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*Journée MissTerre, 19-23 novembre 2012*

# The abrupt 4xCO<sub>2</sub> experiment



$$N \approx \Delta Q_{4x} + \lambda \Delta T$$

(Gregory et al., 2004,  
Gregory and Webb,  
2008)

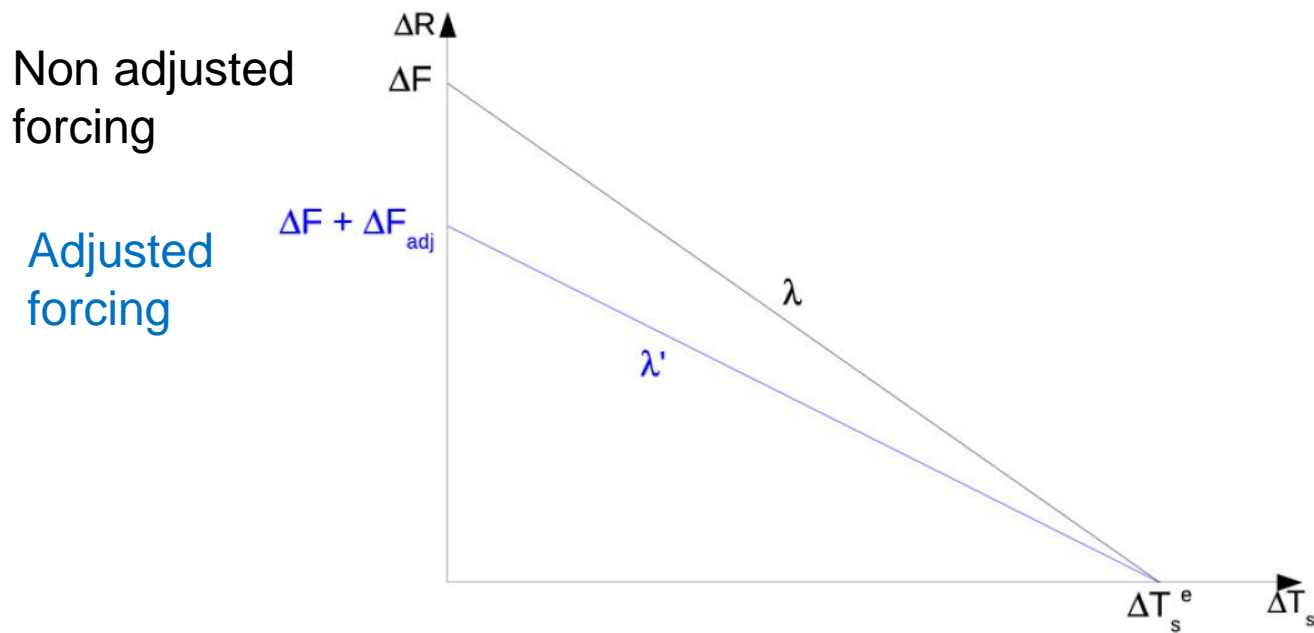
Linear regression between  $N$  and  $\Delta T$

$\Delta Q_{4x}$  : intercept

$\lambda$  : slope

Classical estimate of  $\Delta Q_{4x}$

Difference in forcing estimate: fast response



$$\Delta R(\Delta CO_2, \Delta T_s) = \underbrace{\Delta R(\Delta CO_2, T)}_{\Delta F' = \Delta F + \Delta F_{adj}} + \underbrace{\Delta R(4 \times CO_2, \Delta T_s)}_{\lambda' \Delta T}$$

abrupt\_4xCO2  
 - piControl  
 $\approx$  abrupt\_4xCO2  
 - SSTclim

SSTclim\_4xCO2  
 - SSTclim

abrupt\_4xCO2 –  
 SSTclim\_4xCO2

# Methode

Decomposition of the adjustment to the forcing:

$$\sum_x \Delta F_x = \Delta F_{adj} \qquad \Delta F_x = \frac{\partial R}{\partial x} \Delta x = K_x \Delta x$$

Kernel method

Decomposition of the feedbacks:

$$\lambda = \sum_x \lambda_x = \lambda_{pk} + \lambda_{lr} + \lambda_{wv} + \lambda_{alb} + \lambda_{cl}$$

$$\Delta R_x = \frac{\partial R}{\partial x} \Delta x = K_x \Delta x$$

$$\lambda_x = \frac{\Delta R_x - \Delta F_x}{\Delta T_s}$$

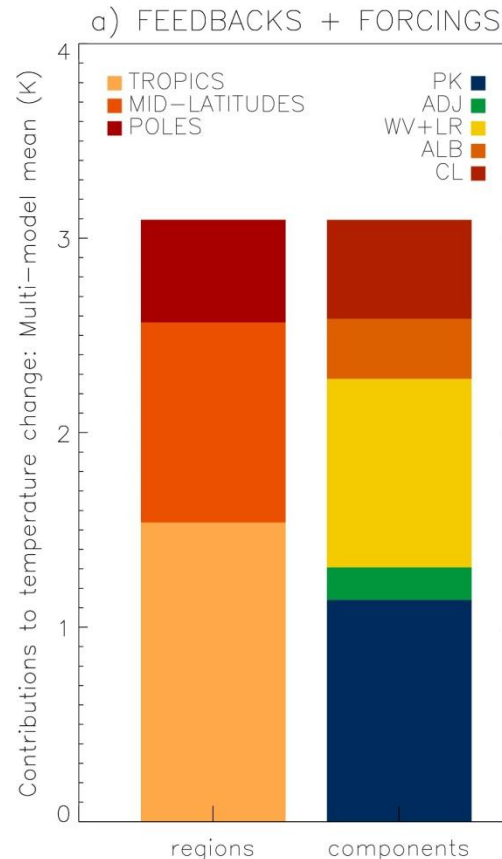
# Decomposition the contribution to temperature increase for a CO2 doubling

$$\Delta T_s = \Delta T_{s,pk} + \Delta T_{s,adj} + \sum_{x \neq pk, adj} \Delta T_{s,x}$$

$$\downarrow \qquad \qquad \downarrow \qquad \qquad \downarrow$$

$$-\frac{\Delta F}{\lambda_{pk}} \qquad -\frac{\sum_x \Delta F_x}{\lambda_{pk}} \qquad -\frac{\lambda_x}{\lambda_{pk}} \Delta T_s$$

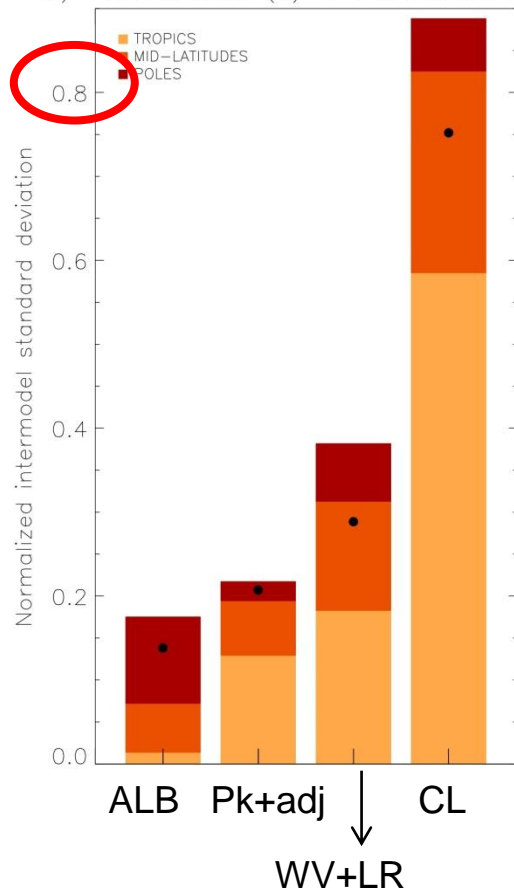
La contribution des  
rétroactions de chacune  
des régions est approx.  
proportionnelle à sa  
surface



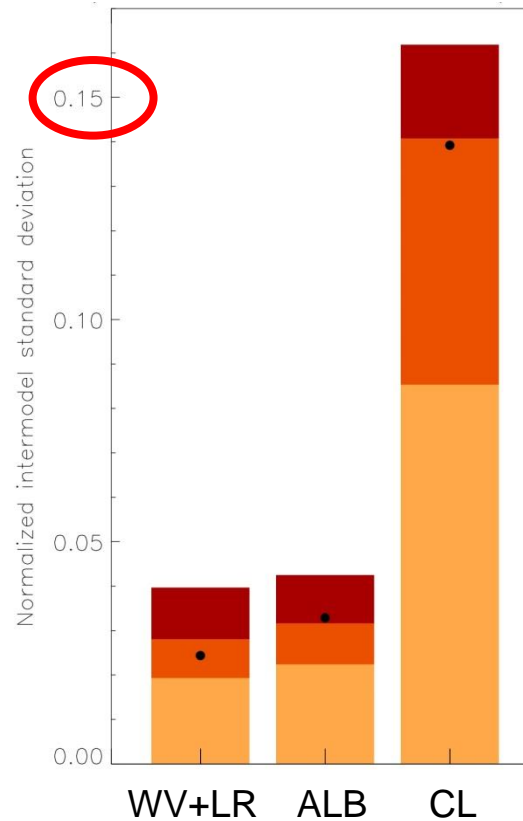
# Decomposition the contribution to the *spread* of the temperature increase for a CO2 doubling

Normalized spread  $\sigma(\Delta T_x) / (\Delta T)$

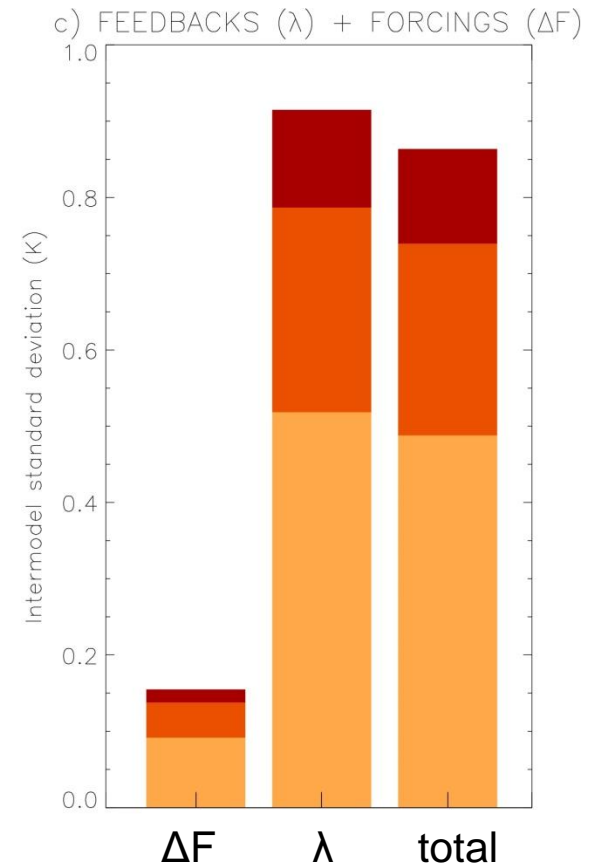
feedback  $\lambda_x$   
+forcing adj  $\Delta F$



forcing adj  $\Delta F_x$

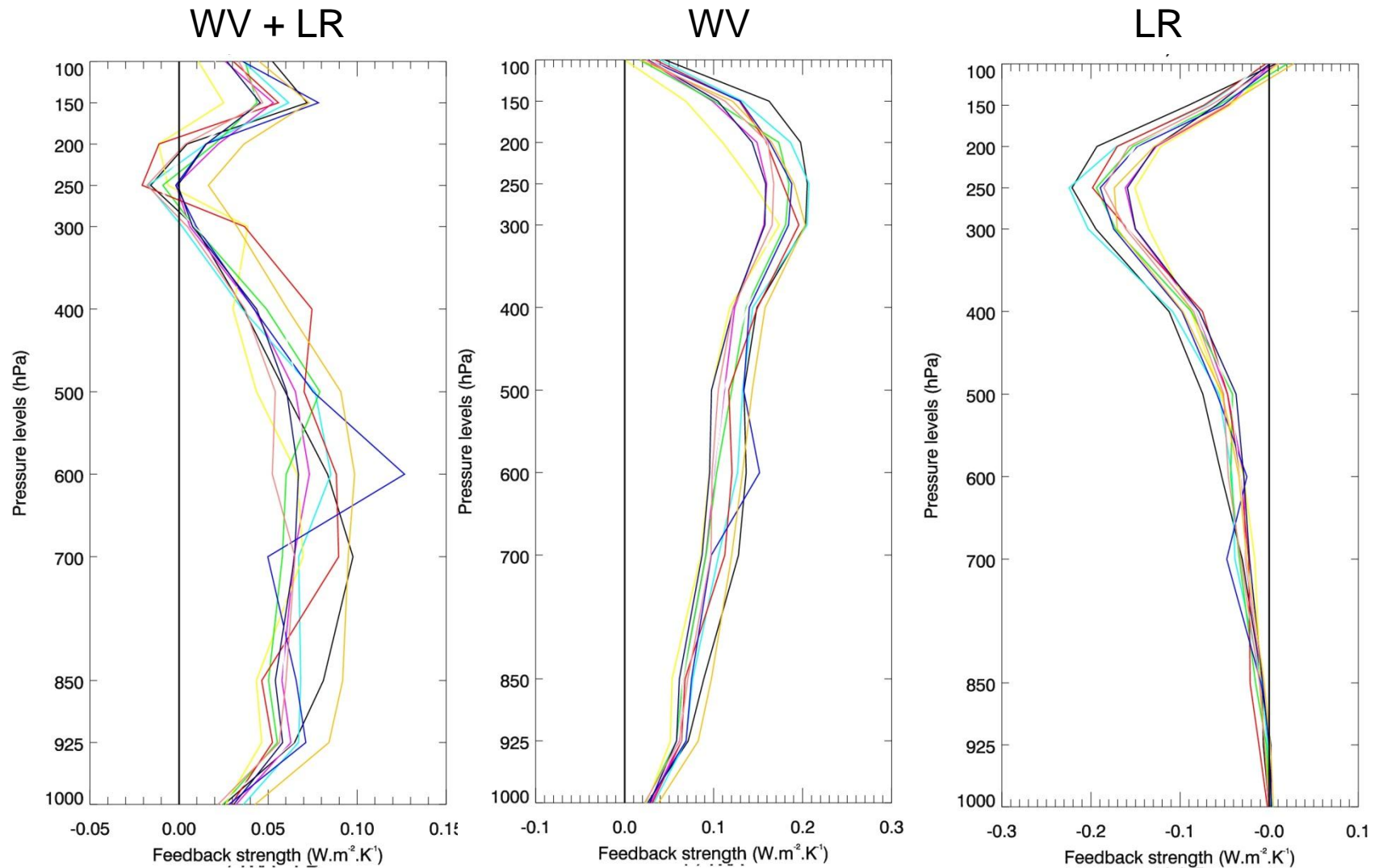


Spread  $\sigma(\Delta T_x)$

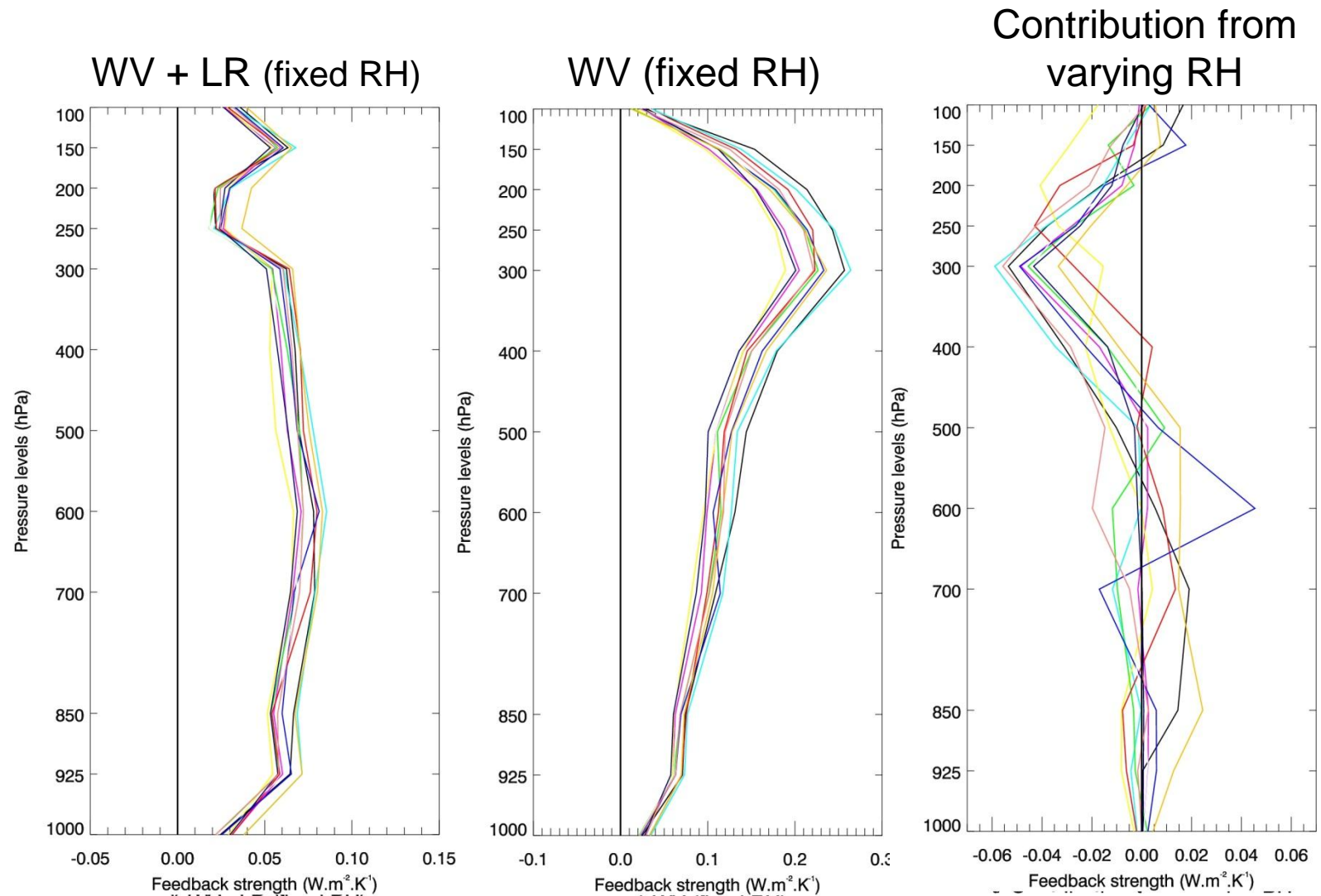


# Water vapor + lapse rate feedback, in the tropics

$$\Delta F_x = \frac{\partial R}{\partial x} \Delta x = K_x \Delta x$$

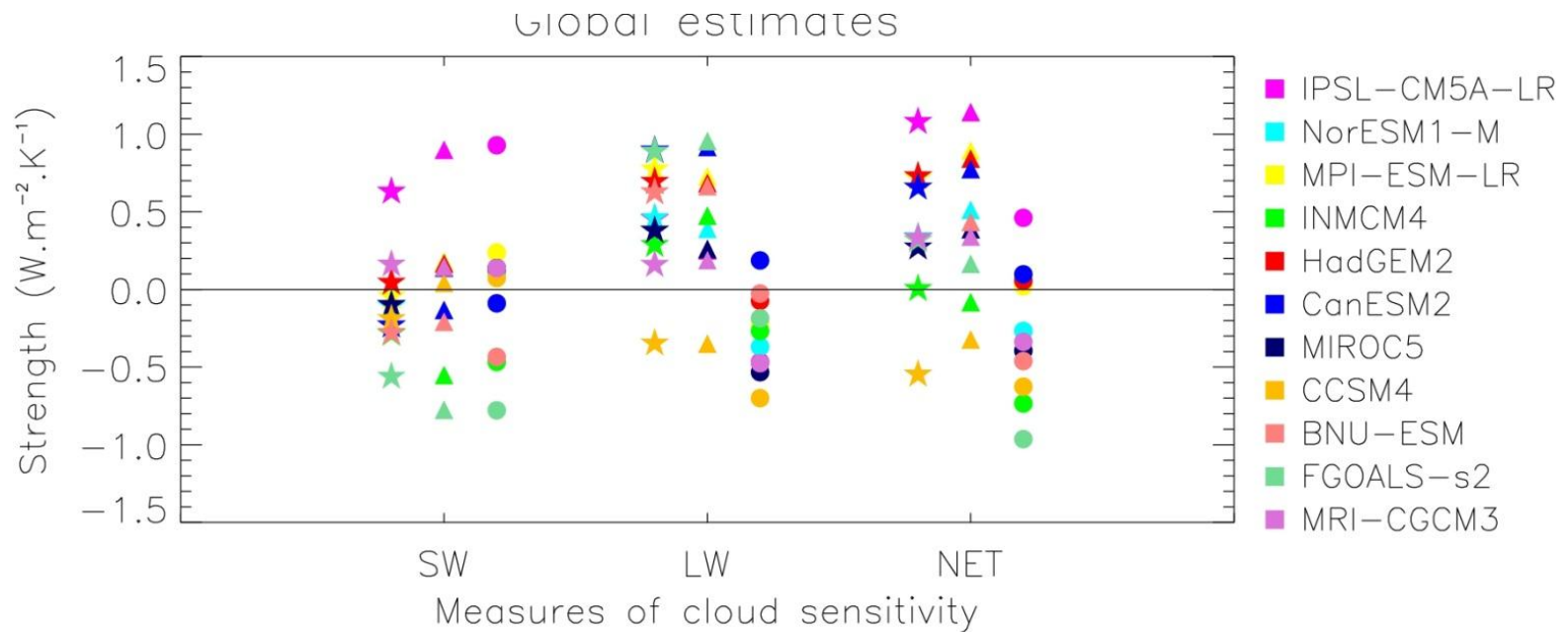


# Water vapor + lapse rate feedback, in the tropics





# Cloud feedback



# Cloud feedback

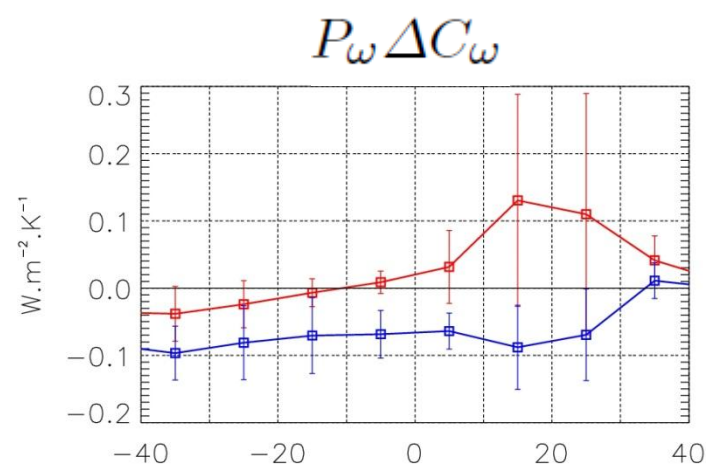
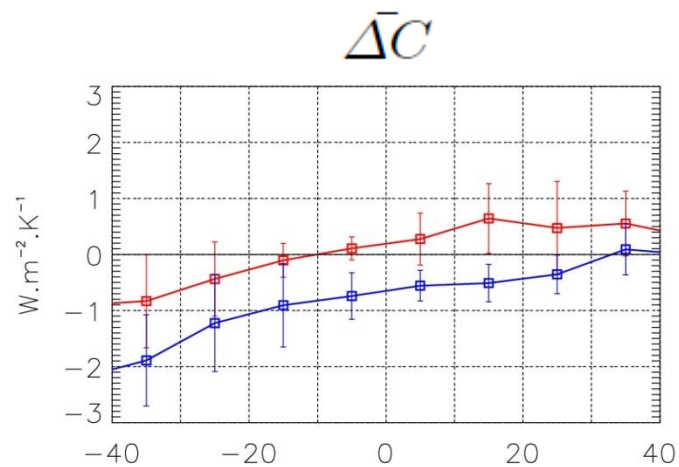
Over the tropical oceans,  
Compositing into different dynamical regimes

$$\bar{C} = \sum_{\omega} P_{\omega} C_{\omega}$$

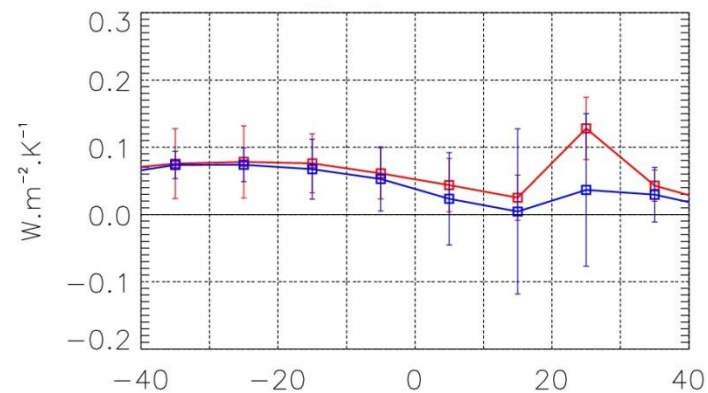
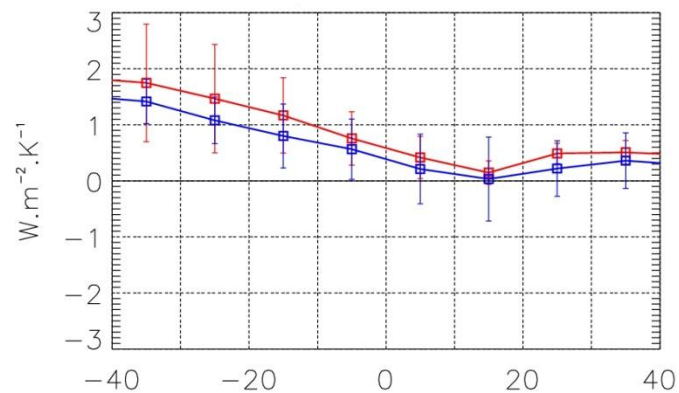
$$\Delta C = \sum_{\omega} C_{\omega} \Delta P_{\omega} + \sum_{\omega} P_{\omega} \Delta C_{\omega} + \sum_{\omega} \Delta C_{\omega} \Delta P_{\omega}$$

Two classes of models: **high sensitive** and **low sensitive** models

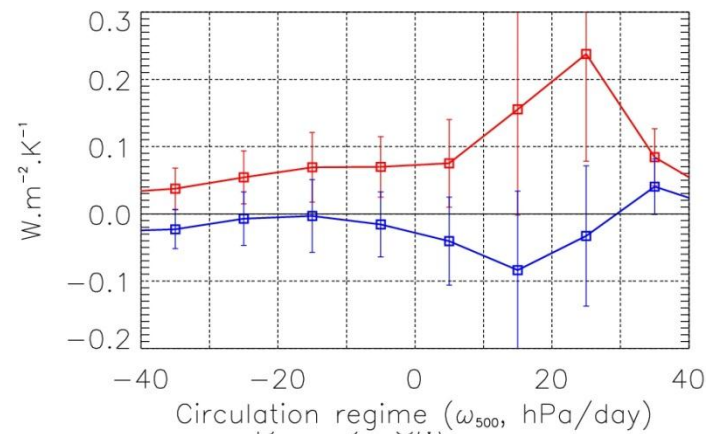
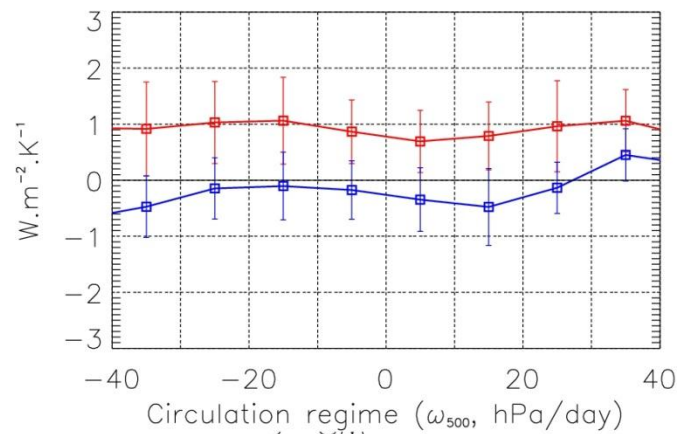
SW



LW



NET



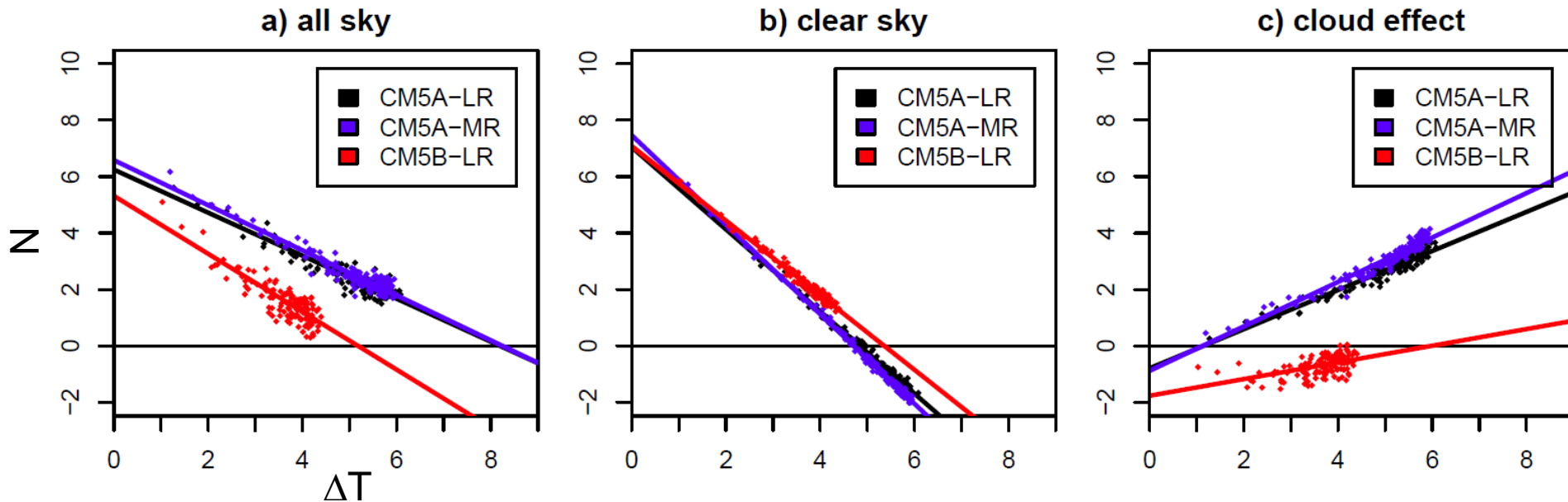
# Conclusion

- Total feedback parameter is reduced by about 10% when considering effect of adjustment on the forcing
- The consideration of adjustment does not reduce the inter-model spread of feedbacks
- Clouds remains the major contributor to the spread of climate sensitivity
- The spread of combined water vapour + lapse rate feedback is entirely due to differences in RH changes
- Spread in tropical clouds: mainly in the SW in region of shallow convection

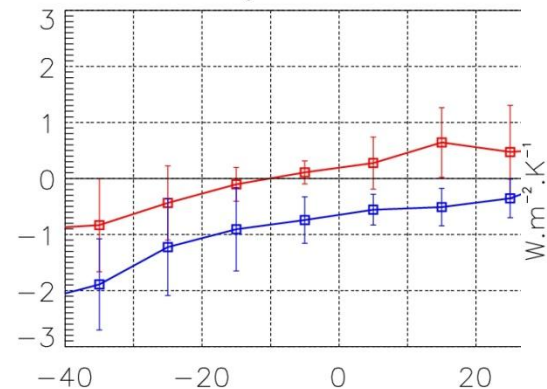
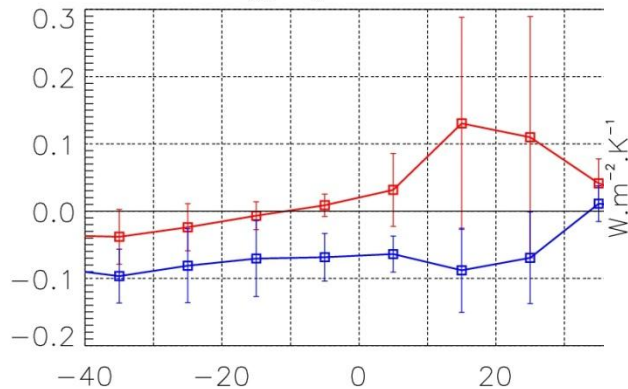
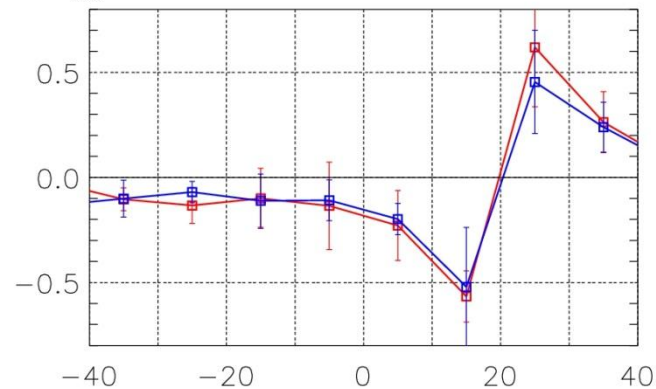
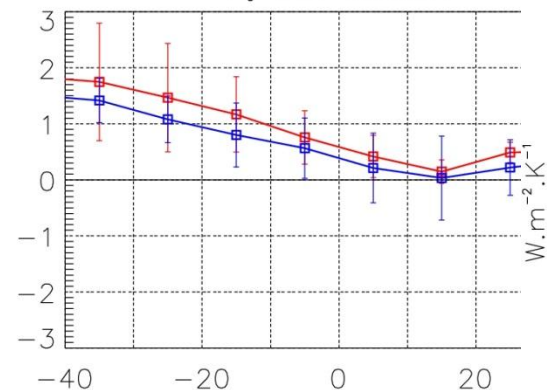
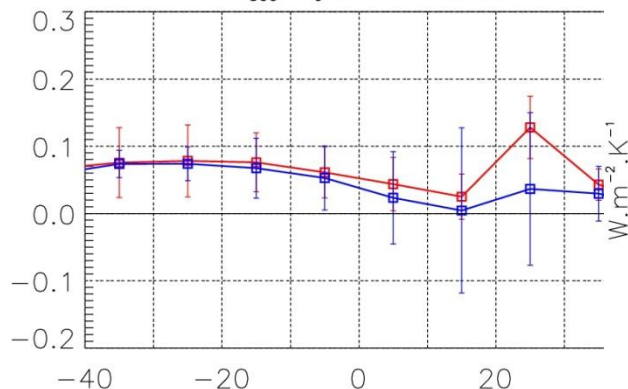
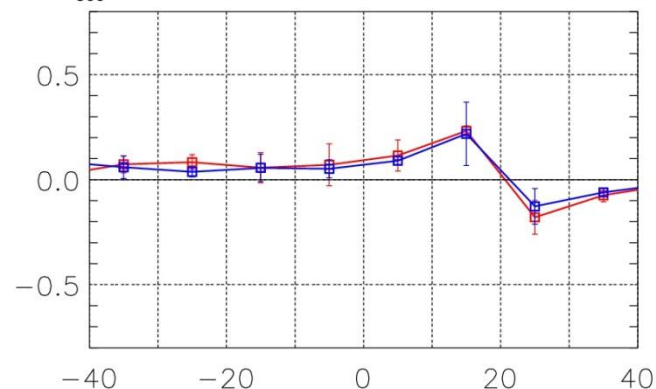
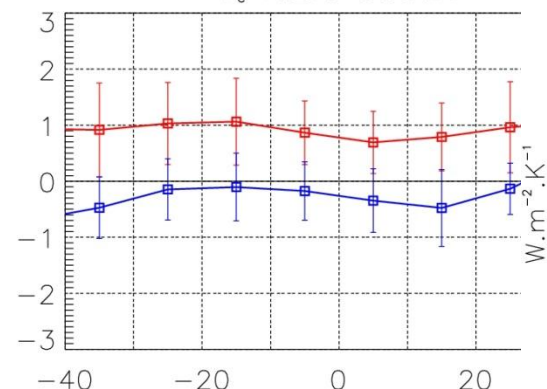
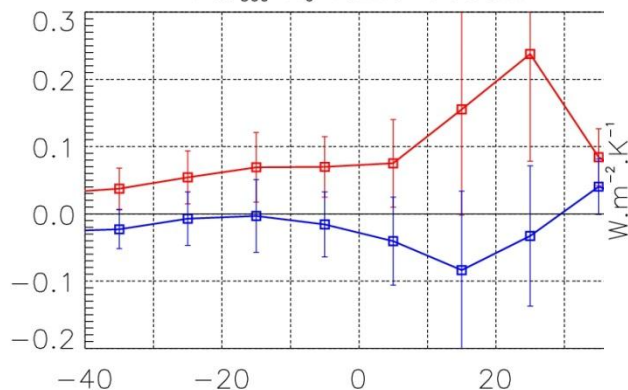
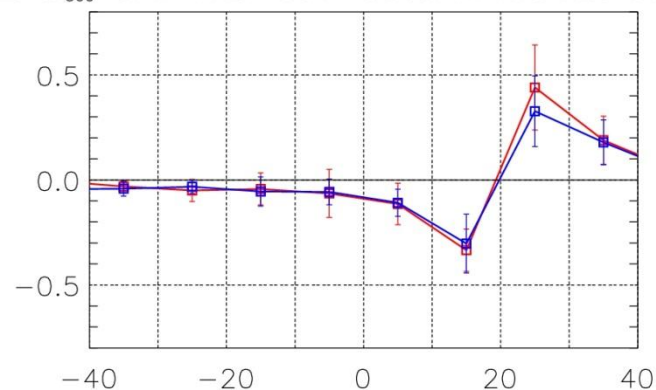


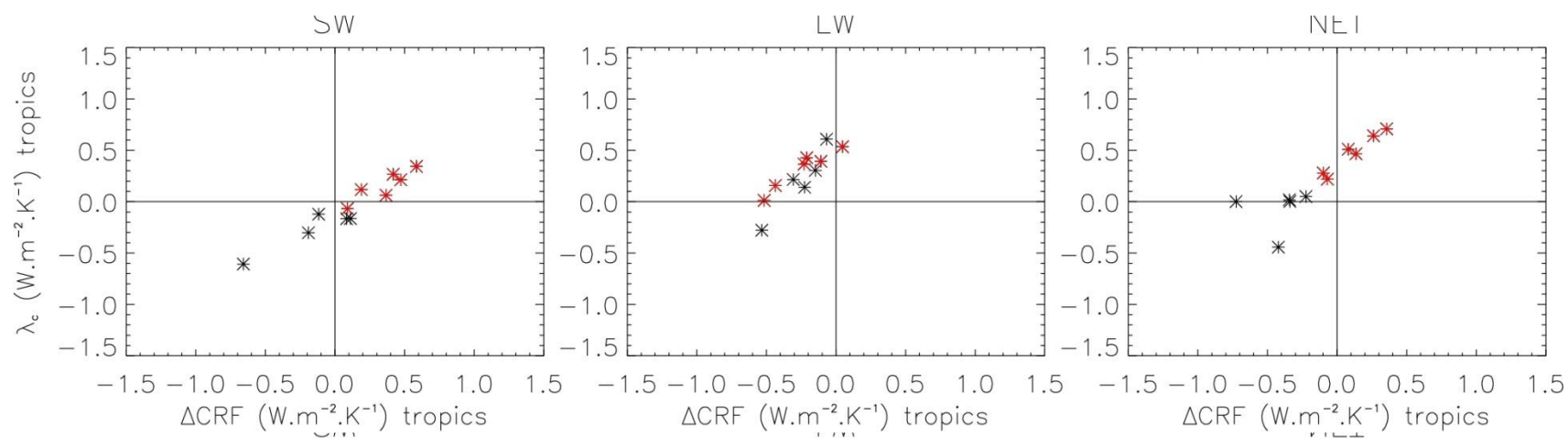
# Climate sensitivity for different IPSL-CM models

Flux TOA vs surface temperature for the abrupt 4xCO<sub>2</sub> simulations

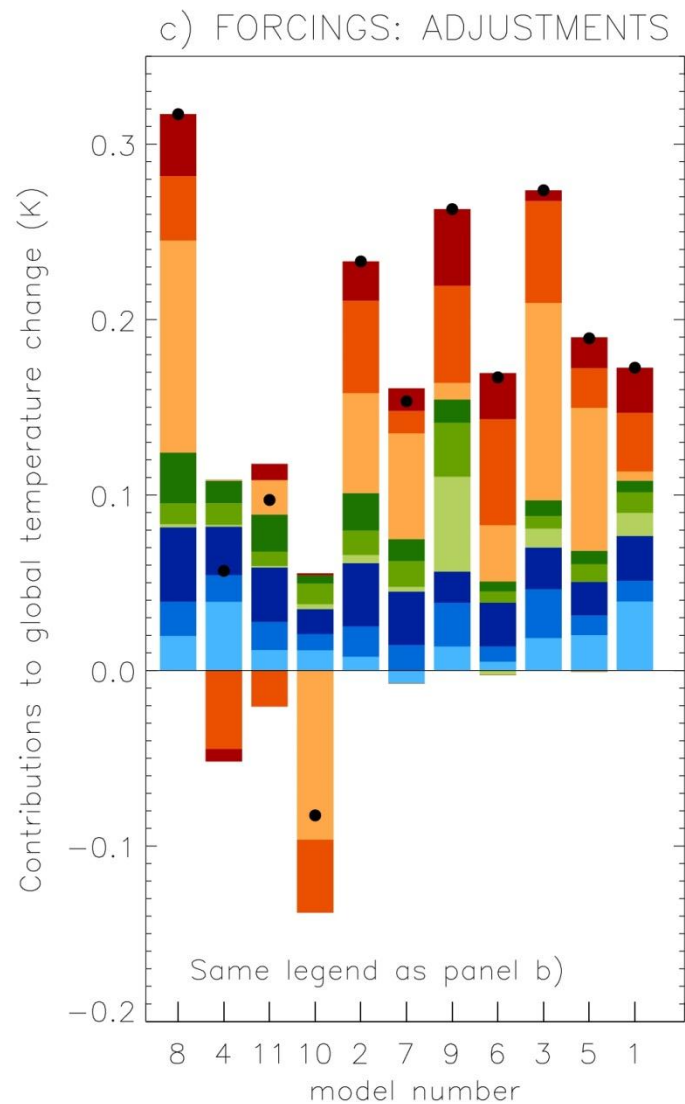
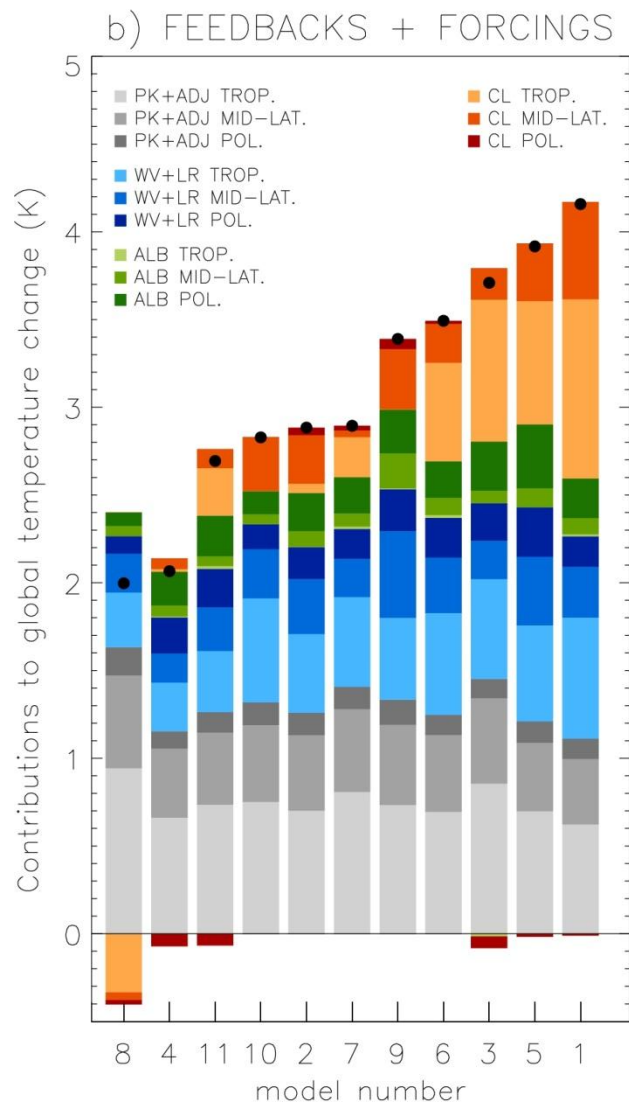


model	1%/year CO <sub>2</sub> increase				abrupt 4xCO <sub>2</sub>		
	$\Delta Q_t(2CO_2)$	$\lambda$	TCR(2CO <sub>2</sub> )	$\Delta T_s^e(2CO_2)$	$\Delta Q_t(2CO_2)$	$\lambda$	$\Delta T_s^e(2CO_2)$
	(Wm <sup>-2</sup> )	(Wm <sup>-2</sup> K <sup>-1</sup> )	(K)	(K)	(Wm <sup>-2</sup> )	(Wm <sup>-2</sup> K <sup>-1</sup> )	(K)
IPSL-CM4	3.5	-0.92	2.13	3.79			
IPSL-CM5A-LR	3.5	-0.98	2.09	3.59	3.12	-0.76	4.10
IPSL-CM5A-MR	3.5	-1.01	2.05	3.47	3.29	-0.80	4.12
IPSL-CM5B-LR	3.5	-1.68	1.52	2.09	2.66	-1.03	2.59

$\lambda_c$  SW: ocean $P\omega_{500} \lambda_c$  SW: ocean $\Delta P\omega_{500}$  CRF SW ctrl-sstClim4xCO2: ocean $\lambda_c$  LW: ocean $P\omega_{500} \lambda_c$  LW: ocean $\Delta P\omega_{500}$  CRF LR ctrl-sstClim4xCO2: ocean $\lambda_c$  NET: ocean $P\omega_{500} \lambda_c$  NET: ocean $\Delta P\omega_{500}$  CRF NET ctrl-sstClim4xCO2: oceanCirculation regime ( $\omega_{500}$ , hPa/day)Circulation regime ( $\omega_{500}$ , hPa/day)Circulation regime ( $\omega_{500}$ , hPa/day)

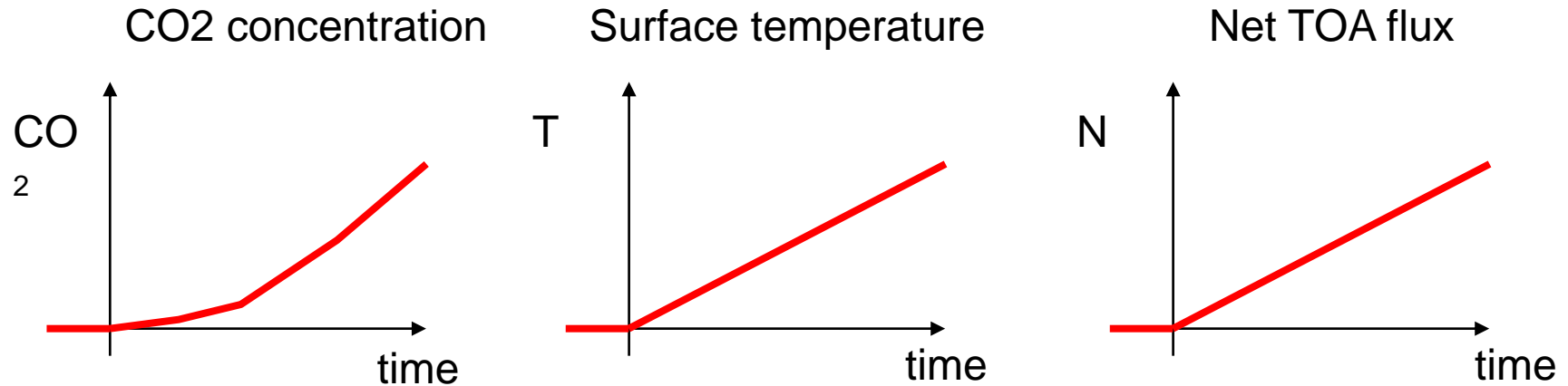






# CMIP5 experiments

## Ramp experiment: 1%/year CO<sub>2</sub> increase



## Step experiment: abrupt 4xCO<sub>2</sub> increase

